Harvard AOSN-II Research: Summary, Recent results and Writing plans

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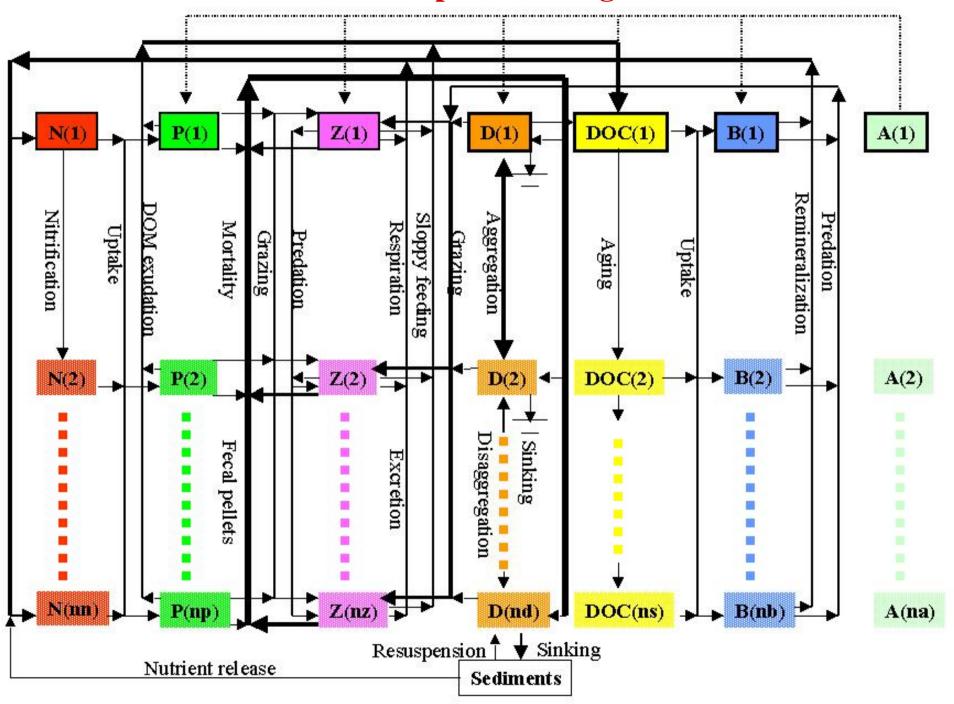
http://www.deas.harvard.edu/~pierrel http://www.deas.harvard.edu/~robinson

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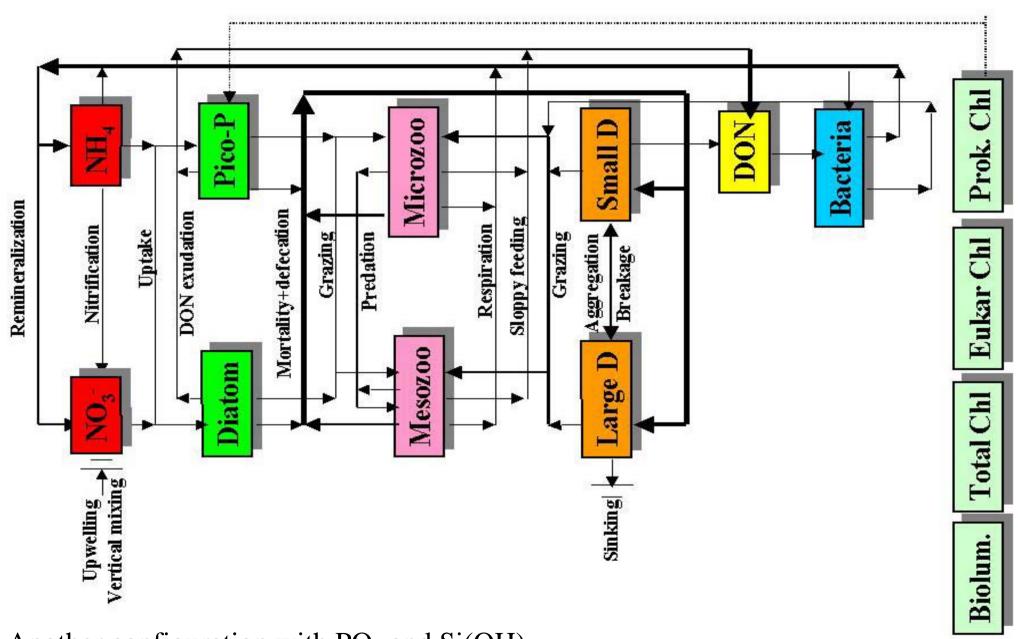


I. Generalized Adaptable Biological Model



(R.C. Tian, P.F.J. Lermusiaux, J.J. McCarthy and A.R. Robinson, HU, 2004)

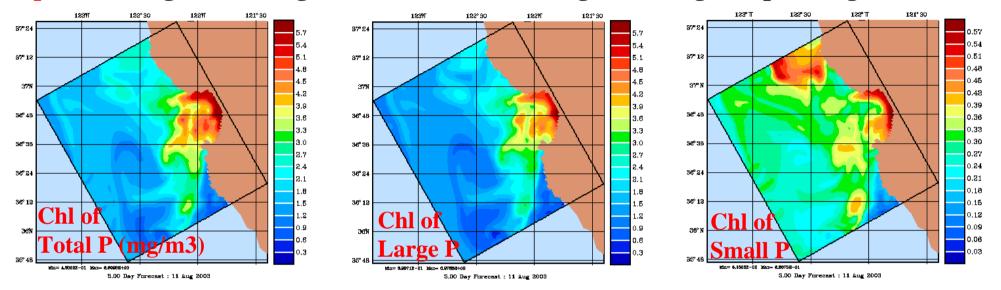
A Priori Biological Model for Monterey Bay



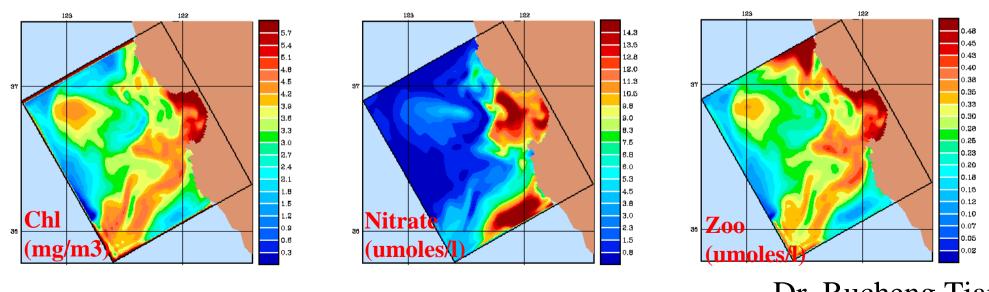
Another configuration with PO₄ and Si(OH)₄

Towards automated quantitative model aggregation and simplification (Older Simulations: done in Sep.-Oct. 2004)

A priori configuration of generalized model on Aug 11 during an upwelling event



NPZ configuration of generalized model on Aug 11 during same upwelling event



Dr. Rucheng Tian

A Priori Biological Modeling Set-Up

1) Field initialization

• Observed variables: Objective analysis of historical-July/Pt Sur data

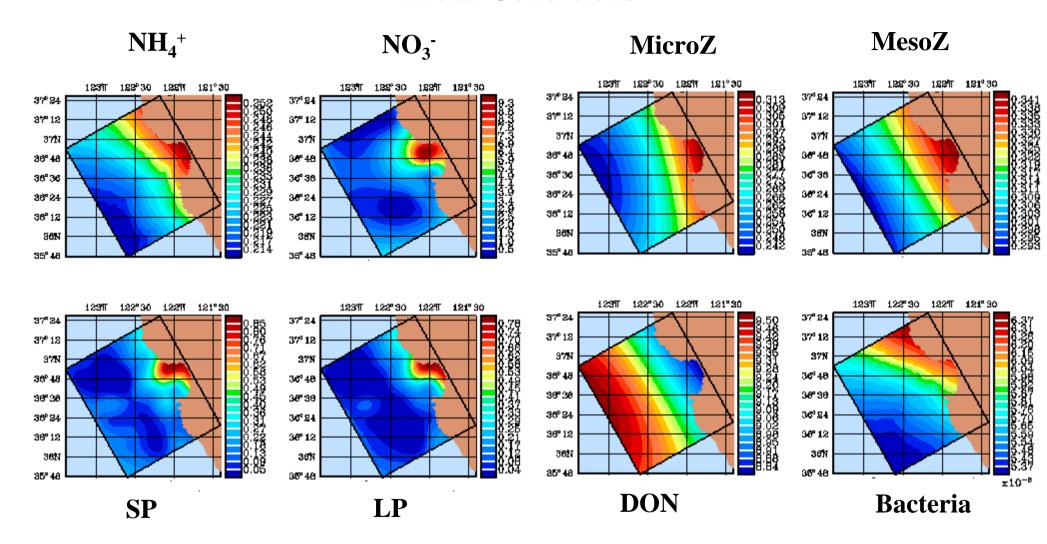
Source	Nb	Parameters	Years	Note
	Profile			
NODC	33	T, S, NO ₃ ,Chlor	1978-1998	
CALCOFI	54	T, S, NO ₃ ,Chlor	1964-1987	
CALCOFI	2	Mesozooplankton (monthly averaged)	1951-1985	<200m
MBARI	30	T, S, NO ₃ ,Chlor(Monthly averaged at discrete depths & years)	1989-2002	<200m
MBARI	3	Microzooplankton (Season averaged)	1989-1998	C1, M1-2
Bauer 1998	3	DOC	1995	35°40'
Bard 1997	3	Bacteria	1993	
Ward 2001	9	Ammonium	1992-1993	Bronk 2001
AOSN II	318	Fluorescence Chlorophyll, Night casts	1-6/08/03	Haddock
AOSN II	98	Nitrate	4-6/08/03	Haddock

- OA correlation scales: 15 km decay, 40 km zero-crossing
- Non-observed variables: Computed from statistical and dynamical (not yet) balances with observed variables (and chosen model)

2) Selection of model parameters and a priori structures/parameterizations

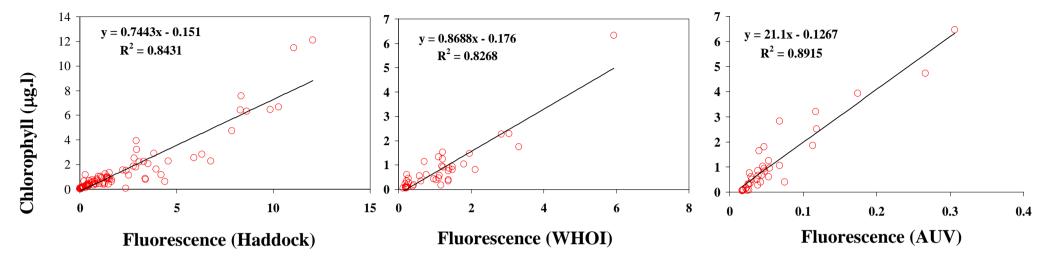
- A priori values chosen as in Olivieri and Chavez (2000) and rest of literature
- Tuning of values and functions by comparing model forecast to AOSN-II data
- 3) Estimation of 1) IC and 2) model parameters/structures not independent
- 4) Assimilation: Only physical data. No biological data assimilation yet

Initial Conditions

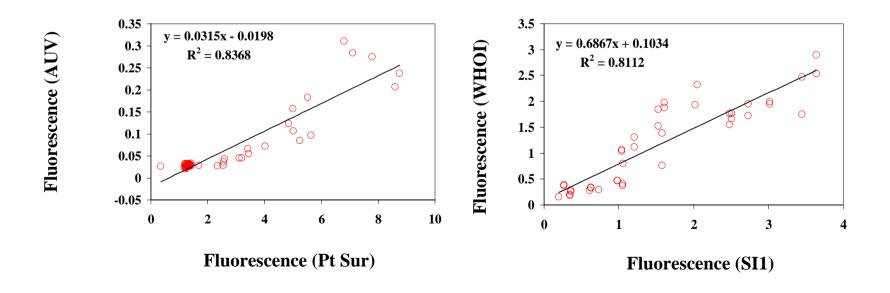


Determined through objective analysis (unit is in m M N m⁻³)

Chlorophyll-Fluorescence relationships



Fluorescence relationships



Cross-Section in Chl µg/l and NO₃:

Observations (S. Haddock et al) vs Simulations

Aug 06 - Aug 18: Upwelling

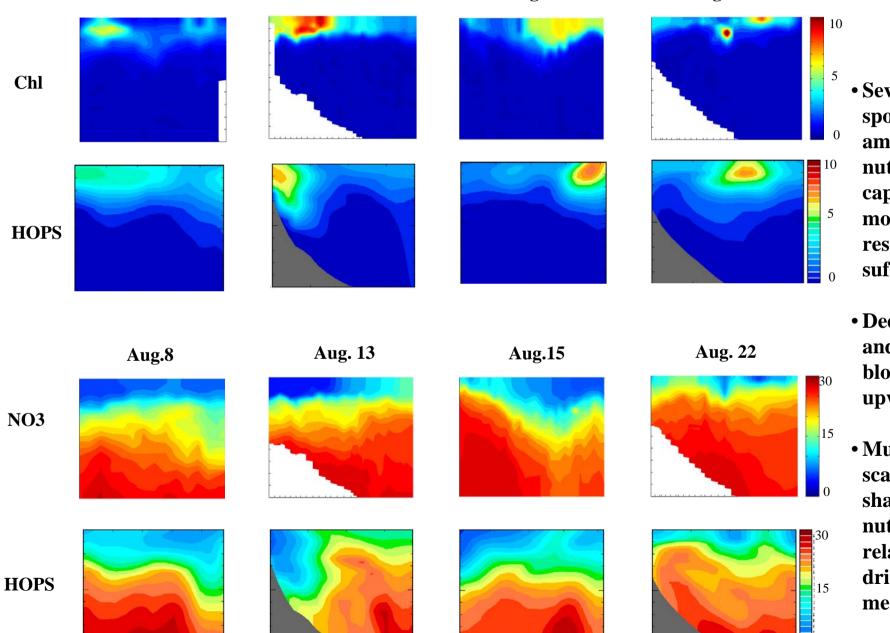
Aug 19 - Aug 23: Relaxation

Aug 27 - Aug 30: Upwelling

Aug. 13

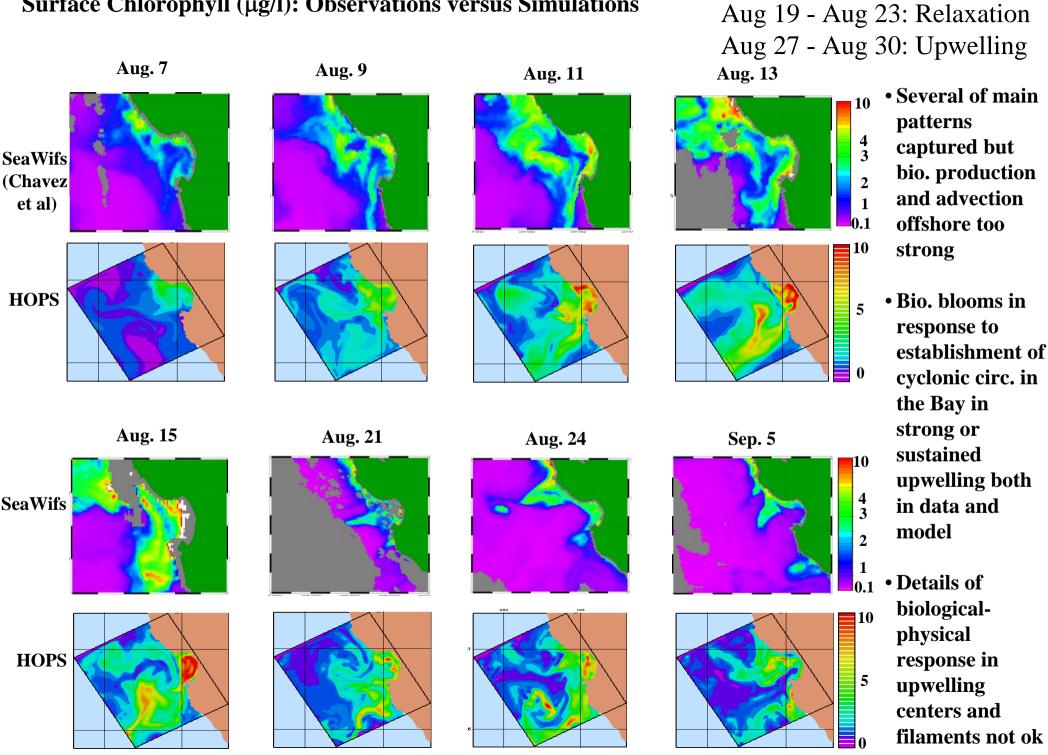
Aug. 15

Aug. 22



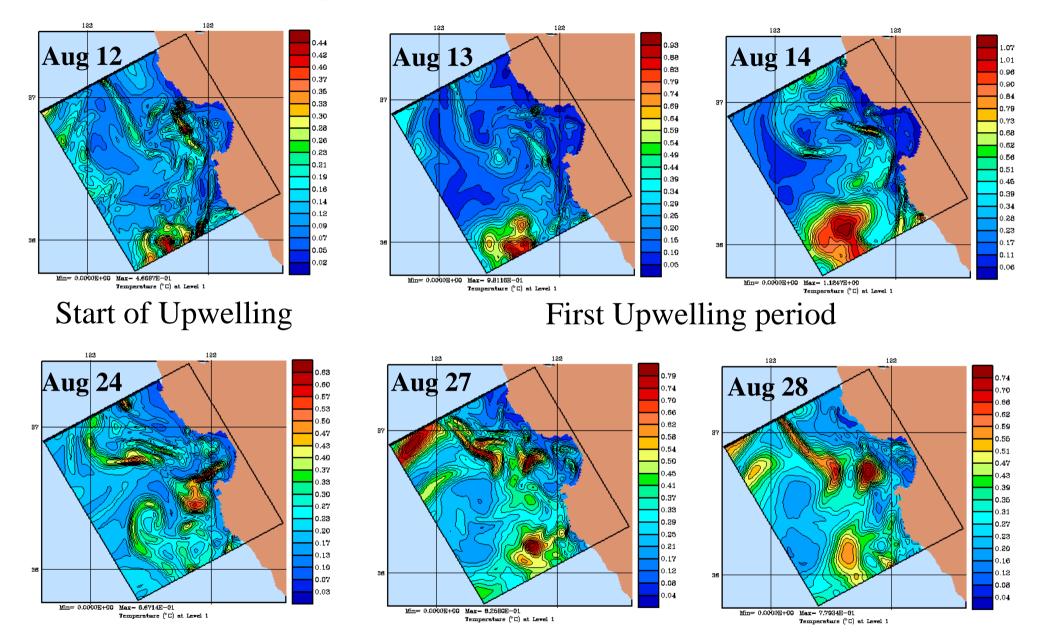
- Several Chl hotspots position and amplitudes, and nutricline tilts, captured but bio. model vertical resolution not sufficient
- Deeper nutricline and stronger blooms during upwelling
- Much smaller scale hot-spots and shallower nutricline during relaxation (oceanic driven submesoscales)

Surface Chlorophyll (µg/l): Observations versus Simulations



Aug 06 - Aug 18: Upwelling

ESSE Surface Temperature Error Standard Deviation Forecasts



End of Relaxation

Second Upwelling period

- •Real-time consistent error forecasting, data assimilation and adaptive sampling (1 month)
- •ESSE results described in details and posted on the Web daily (see Wayne's AOSN2 page)

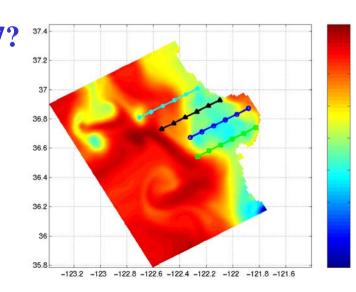
II.a Adaptive sampling via ESSE: evolving predicted objective field

- Objective: Minimize predicted trace of full error covariance (T,S,U,V error std Dev).
- Scales: Strategic/Experiment (not tactical yet). Day to week.
- Assumptions: Small number of pre-selected tracks/regions (based on quick look on error forecast and constrained by operation)
- Problem solved: e.g. Compute today, the tracks/regions to sample tomorrow, that will most reduce uncertainties the day after tomorrow.
- Objective field changes during computation and is affected by data to-be-collected
- Model errors Q can account for coverage term

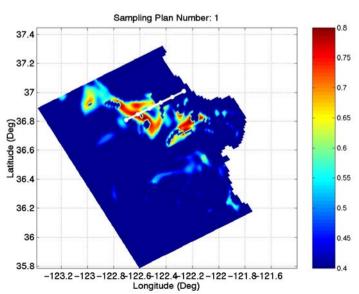
Example: Which sampling on Aug 26 optimally reduces uncertainties on Aug 27?

- Based on nonlinear error covariance evolution
- For every choice of adaptive strategy, an ensemble is computed

4 candidate tracks, overlaid on surface T fct for Aug 26



Best predicted relative error reduction: track 1



II.b Optimal Paths Generation for a "fixed" objective field

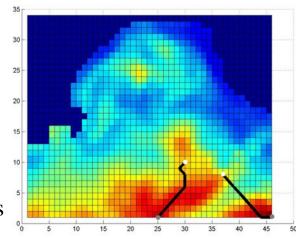
(Namik K. Yilmaz, P. Lermusiaux and N. Patrikalakis)

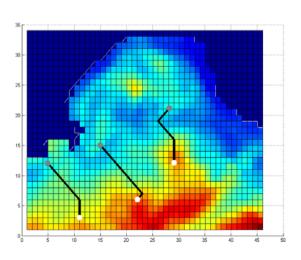
- Objective: Minimize ESSE error standard deviation of temperature field
- Scales: Strategic/Tactical
- Assumptions
 - Speed of platforms >> time-rate of change of environment
 - Objective field fixed during the computation of the path and is not affected by new data
- Problem solved: assuming the error is like that now and will remain so for the next few hours, where do I send my gliders/AUVs?
- Method: Combinatorial optimization (Mixed-Integer Programming, using Xpress-MP code)
 - Objective field (error stand. dev.) represented as a piecewise-linear: solved exactly by MIP
 - Possible paths defined on discrete grid: set of possible path is thus finite (but large)
 - Constraints imposed on vehicle displacements dx, dy, dz for meaningful path

Example for Two and Three Vehicles, 2D objective field

Grey dots: starting points

White dots: MIP optimal end points





III. Progress towards near-inertial and tidal modeling for AOSN-II

•Free surface HOPS

-Currently calibrated for MREA-03 (Elba-Corsican Channel) and MREA-04 (Portuguese coastal waters)

•Hierarchy of tidal parameterizations

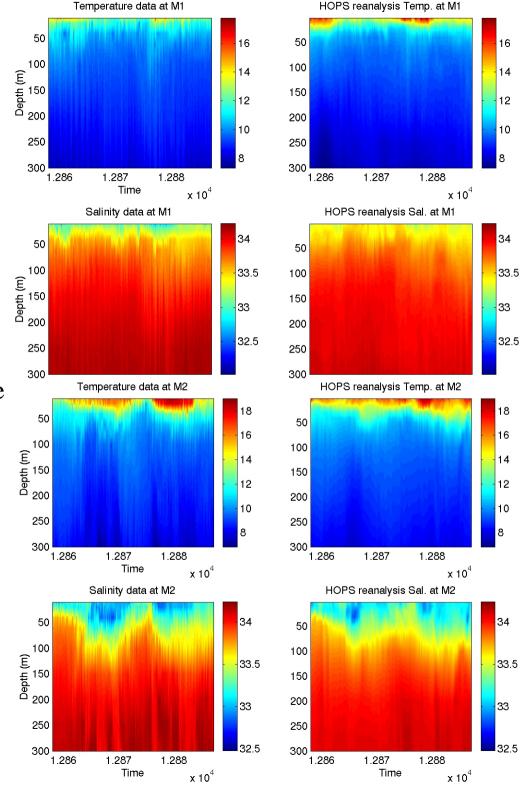
- -Rigid-lid HOPS:
 - Reynolds stresses (vertical, horizontal)
 - Horiz. tidal advection of tracers (1/2 free surface)
- -Free-surface HOPS
- -ESSE stochastic forcings (estimate parameters from data)

•Evaluation plans

- -Compare models and data at M1/M2
- Compare models and glider data

•Initiate research towards

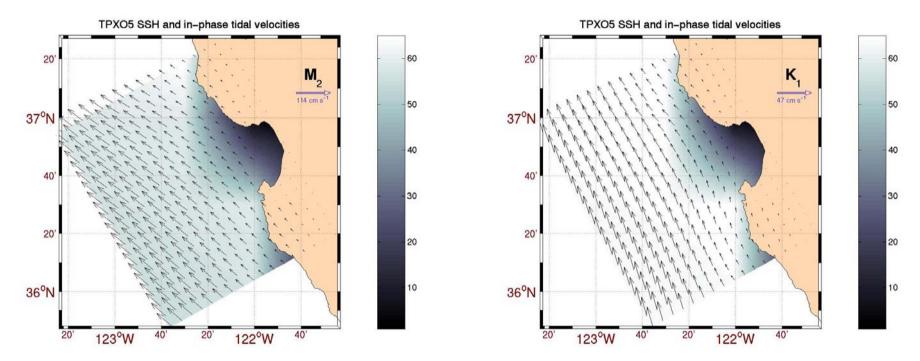
Optimal gliders patterns for sampling/filtering missing scales



HOPS T. power spectral dens. at M1 (1 to 10 m) Measured T. power spectral dens. at M1 (10 m) 10⁰ 10⁰ **10**⁻² **10**⁻² 10 m Diurnal 1 m Inertial 5 m 10 m Semi-diurnal **10**⁻² **10**⁻² 10⁻¹ 10⁻¹ HOPS T. power spectral dens. at M1 (15 to 300 m) Measured T. power spectral dens. at M1 (20 to 300 m) 10⁰ 10⁰ Diurnal Diurnal 10^{-5} 10⁻⁵ Inertial Inertial Semi-diurnal Semi-diurnal 10⁻² 10⁻² 10⁻¹ 10⁻¹

Modeling of tidal effects in HOPS

- Obtain first estimate of principal tidal constituents via a shallow water model
 - 1. Global TPXO5 fields (Egbert, Bennett et al.)
 - 2. Nested regional OTIS inversion using tidal-gauges and TPX05 at open-boundary



Need to cross-evaluate with Igor Shulman and Leslie Rosenfeld

- Used to estimate hierarchy of tidal parameterizations :
 - Vertical tidal Reynolds stresses (diff., visc.) $K_T = \alpha / |u_T|/2$ and $K = max(K_S, K_T)$
 - ii. Modification of bottom stress
 - iii. Horiz. momentum tidal Reyn. stresses
 - iv. Horiz. tidal advection of tracers
 - v. Forcing for free-surface HOPS

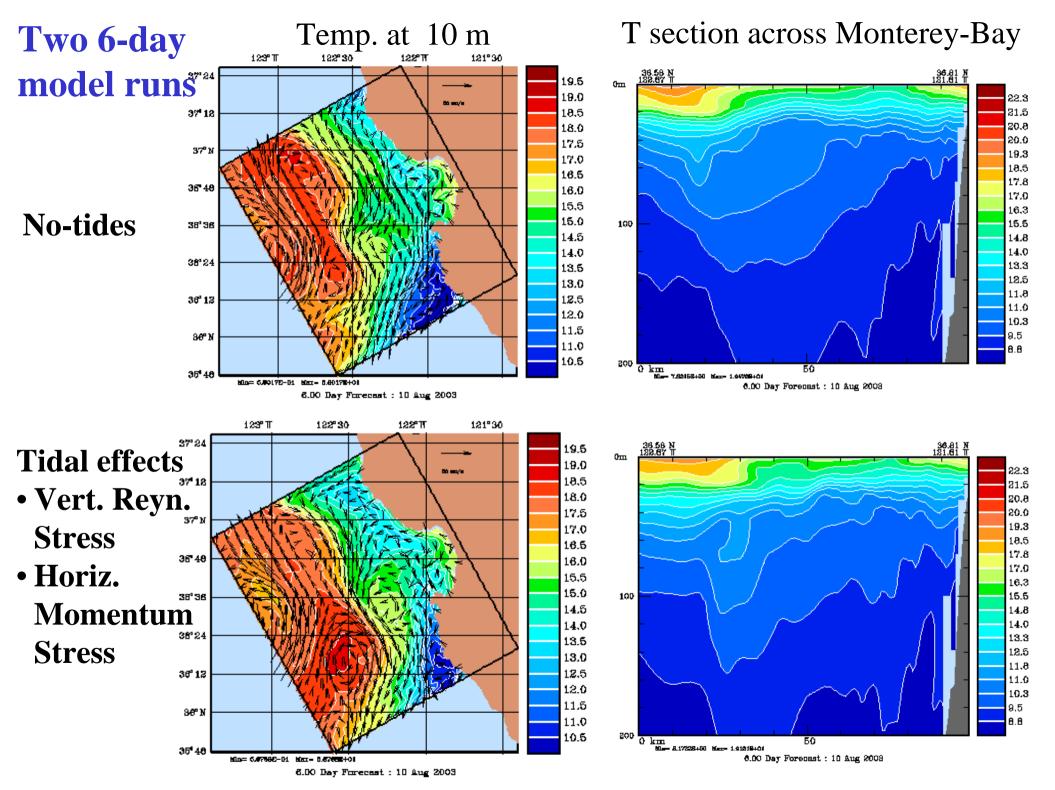
$$K_T = \alpha / |\boldsymbol{u}_T|/2$$
 and $K = max(K_S, K_T)$

$$\tau = C_D / |u_{S+} u_T| / |u_S|$$

 Σ_{ω} (Reyn. stresses averaged over own T_{ω})

½ free surface

full free surface



IV. Flux Balances and Term-by-term Balances

• Physical model: Primitive-Equation (PDE, x, y, z, t: HOPS)

Horiz. Mom.
$$\frac{D\mathbf{u_h}}{Dt} + f \, \mathbf{e}_3 \wedge \mathbf{u}_h = -\frac{1}{\rho_0} \nabla_h \, p_w + \nabla_h \cdot (A_h \nabla_h \mathbf{u}_h) + \frac{\partial A_v \, \partial \mathbf{u}_h / \partial z}{\partial z}$$

Vert. Mom.
$$\rho g + \frac{\partial p_w}{\partial z} = 0$$

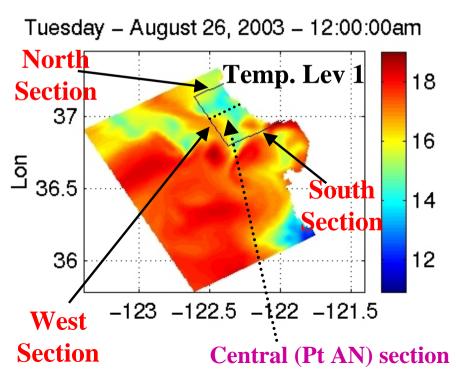
Thermal en.
$$\frac{DT}{Dt} = \nabla_h \cdot (K_h \nabla_h T) + \frac{\partial K_v \partial T/\partial z}{\partial z}$$

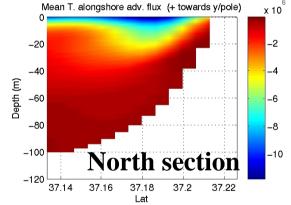
Cons. of salt
$$\frac{DS}{Dt} = \nabla_h \cdot (K_h \nabla_h S) + \frac{\partial K_v \partial S/\partial z}{\partial z}$$

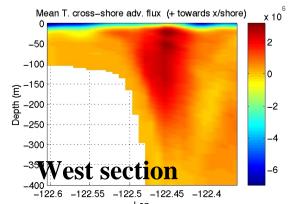
Cons. of mass $\nabla \cdot \mathbf{u} = 0$

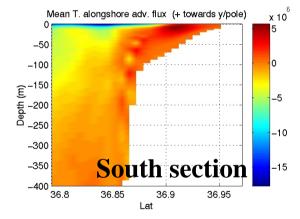
Eqn. of state $\rho(\mathbf{r},z,t) = \rho(T,S,p_w)$

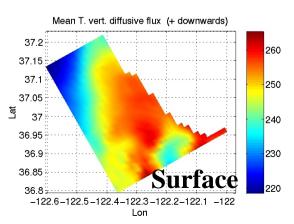
Heat Flux Balances: 4 fluxes normal to each side averaged over first upwelling period Mean Fluxes (W/m2) over: August 6, 2003 - 10:30:00pm -> August 13, 2003 - 4:30:00am GMT



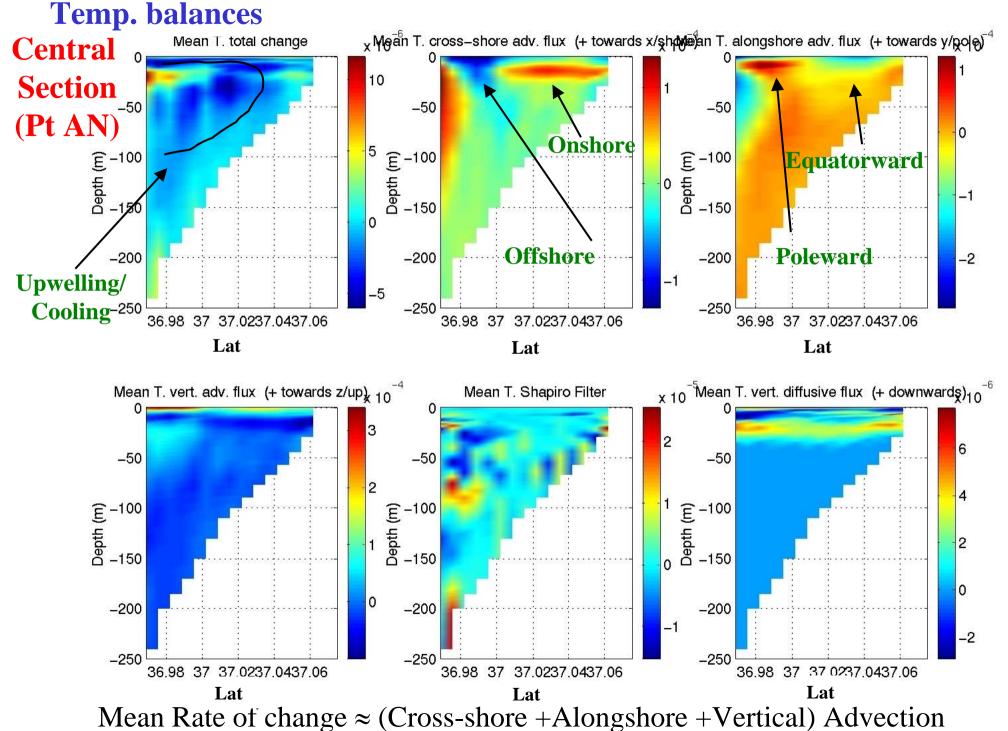






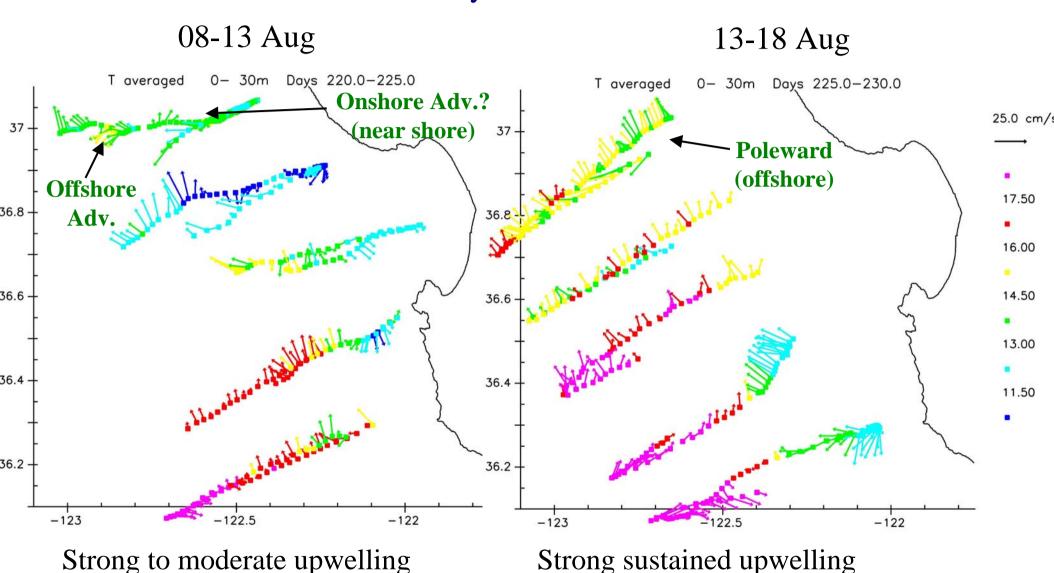


Mean Term-by-Term over: August 6, 2003 – 10:30:00pm -> August 17, 2003 – 1:30:00am GMT

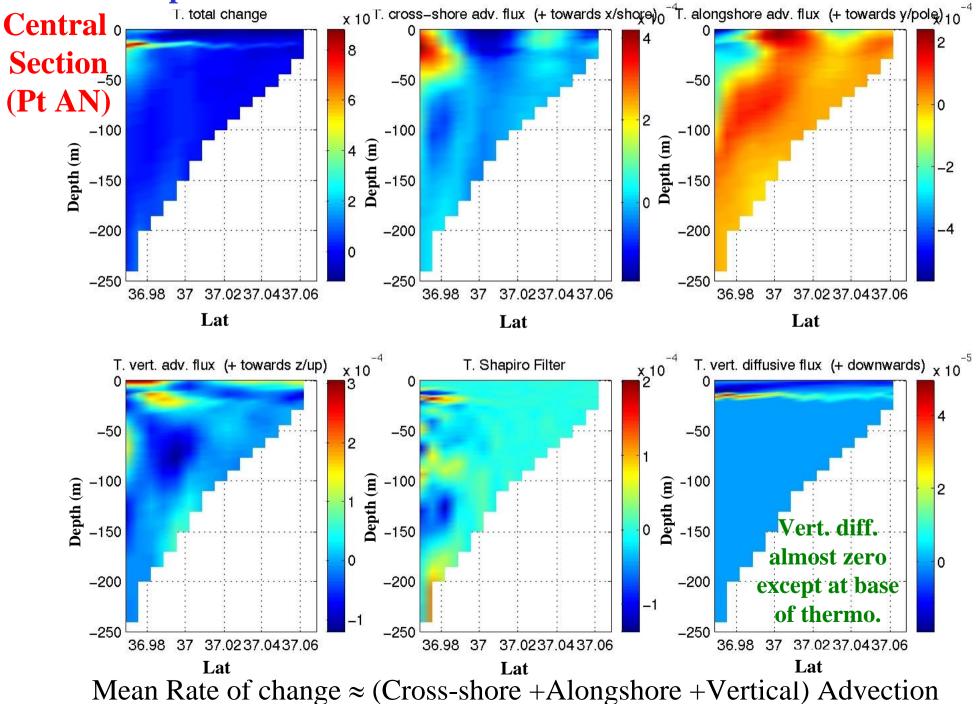


What the Data Shows (Russ Davis et al.)

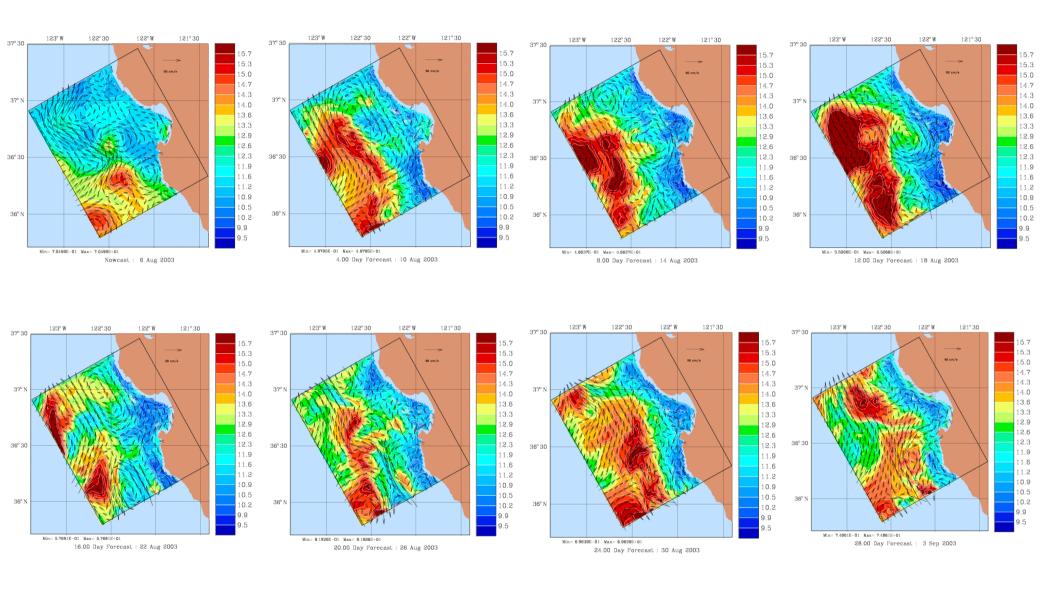
Days 220-230

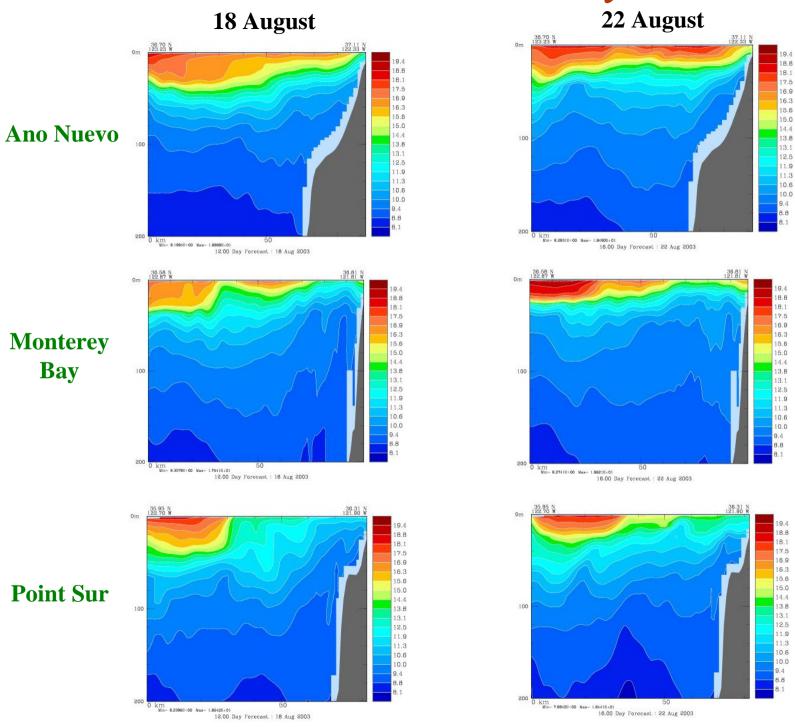


Temp. balances



30m Temperature: 6 August – 3 September (4 day intervals)





The reanalysis simulation is a 35 day simulation spanning 0000Z August 6, 2003 through 0000Z September 10, 2003. The simulation is initialized with and/or assimilates Pt. Sur, John Martin and Pt. Lobos CTDs; WHOI and SIO gliders; and NPS aircraft SST for the period Aug. 2 - Sept. 6, 2003. The model is forced with wind stress, net heat flux and E-P derived from the hourly 3km COAMPS fields.

Two data files are available containing the gzipped netcdf formatted fields. Both files are from the same re-analysis PE run and are on the HOPS horizontal grid.

File 1 has output on the flat levels: 0, -5, -10, -15, -20, -30, -40, -50, -60, -75, -100, -125, -150, -200, -250, -300 with an output frequency of 90 minutes, spanning 0000Z August 6, 2003 through 0000Z September 7, 2003.

File 2 has output on the flat levels: -1, -5, -10, -15, -20, -30, -40, -50, -60, -75, -100, -125, -150, -200, -250, -300, -400, -500, -600, -700, -800, -900, with an output frequency of 3 hours, spanning 0000Z August 6, 2003 through 0000Z September 10, 2003.

Initialization/Boundary Conditions

The initialization field was created by an objective analysis of the August 2-6, 2003 Point Sur CTD + WHOI and SIO glider + NPS SST data (thinned for speed).

The fields were created with the following correlation scales:

XZERO = YZERO = 37.5 (km);

XDCAY = YDCAY = 15.0 (km);

TDCAY = 10.0 (days);

XZEROM = YZEROM = 112.5 (km)

XDCAYM = YDCAYM = 45.0 (km)

TDCAYM = 1000.0 (days)

OADAY = 12858.0 (06 Aug. 2003 - 0:00:00)

Time evolving boundary values were obtained by objective analysis of data concurrent with the remaining Pt. Sur Surveys.

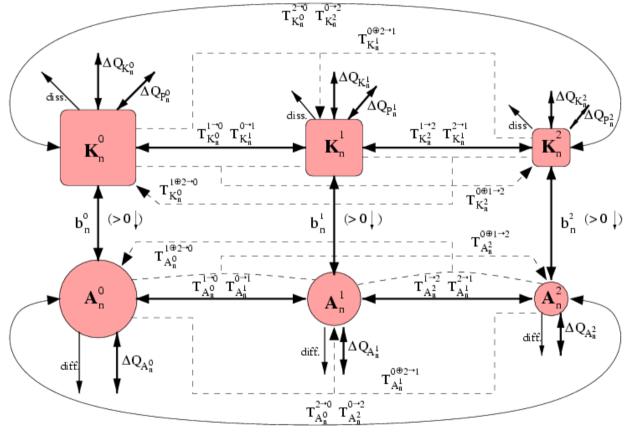
http://people.deas.harvard.edu/~leslie/AOSNII/REANALYSIS/index.html

MS-EVA is a new methodology utilizing multiple scale window decomposition in space and time for the investigation of processes which are:

- multi-scale interactive
- nonlinear
- intermittent in space
- episodic in time

Through exploring:

- pattern generation and
- energy and enstrophy
 - transfers
 - transports, and
 - conversions



MS-EVA helps unravel the intricate relationships between events on different scales and locations in phase and physical space.

Dr. X. San Liang

Window-Window Interactions: MS-EVA-based Localized Instability Theory

Perfect transfer:

A process that exchanges energy among distinct scale windows which does not create nor destroy energy as a whole.

In the MS-EVA framework, the perfect transfers are represented as field-like variables. They are of particular use for real ocean processes which in nature are non-linear and intermittent in space and time.

Localized instability theory:

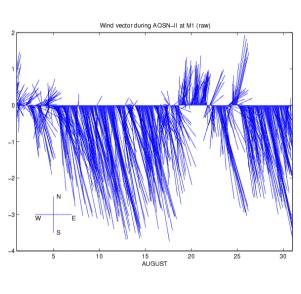
BC: Total perfect transfer of APE from large-scale window to meso-scale window.

BT: Total perfect transfer of KE from large-scale window to meso-scale window.

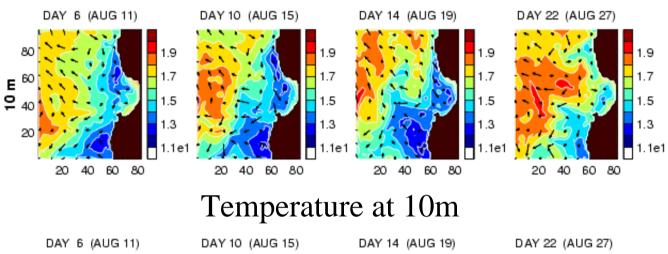
BT + BC > 0 => system locally unstable; otherwise stable

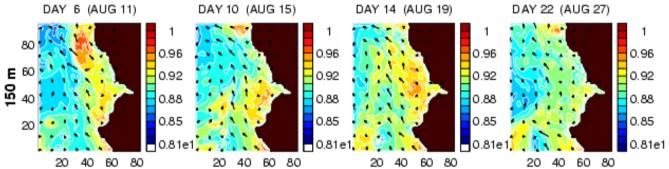
If BT + BC > 0, and

- $BC \le 0 \Rightarrow$ barotropic instability;
- $BT \le 0 \Rightarrow$ baroclinic instability;
- BT > 0 and BC > 0 => mixed instability



M1 Winds

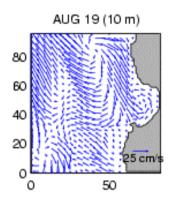




Temperature at 150m

Multi-Scale Window Decomposition in AOSN-II Reanalysis





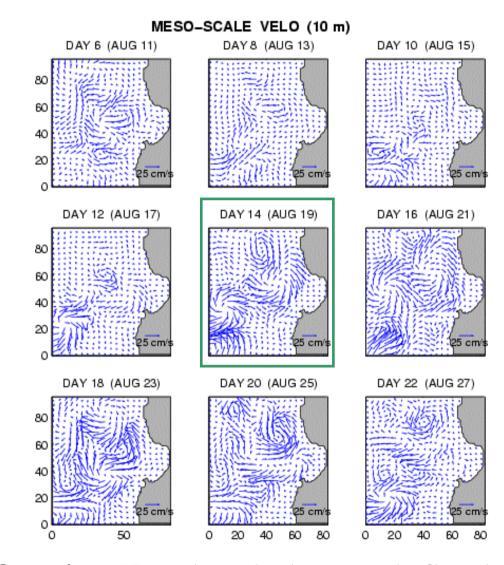
The reconstructed largescale and meso-scale fields are filtered in the horizontal with features < 5km removed.

Time windows

Large scale: > 8 days

Meso-scale: 0.5-8 days

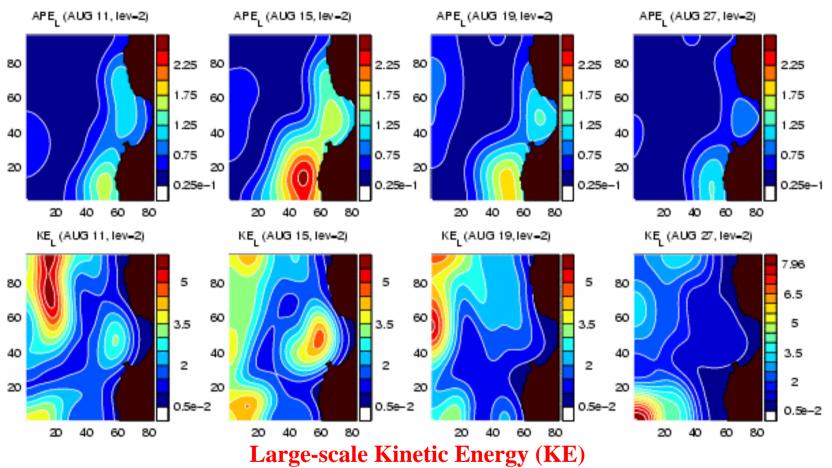
Sub-mesoscale: < 0.5 day



Question: How does the large-scale flow lose stability to generate the meso-scale structures?

• Decomposition in space and time (wavelet-based) of energy/vorticity eqns.



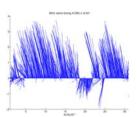


- Both APE and KE decrease during the relaxation period
- Transfer from large-scale window to mesoscale window occurs to account for decrease in large-scale energies (as confirmed by transfer and mesoscale terms)

Windows: Large-scale (>= 8days; > 30km), mesoscale (0.5-8 days), and sub-mesoscale (< 0.5 days)

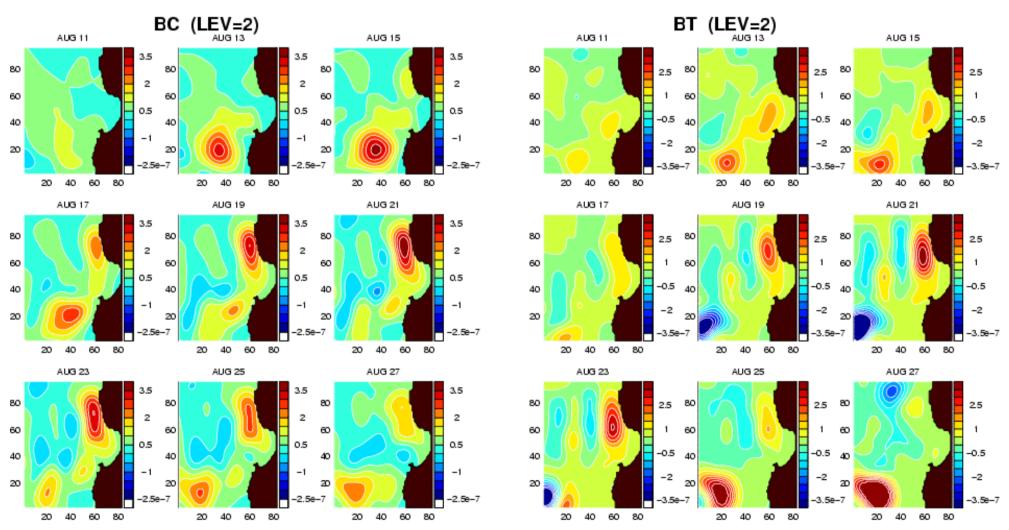
Dr. X. San Liang

MS-EVA Analysis: 11-27 August 2003



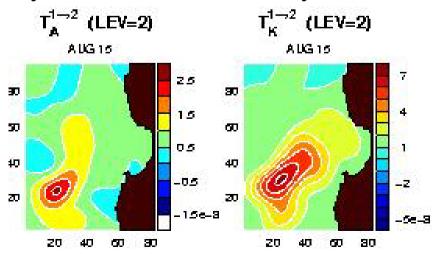
Transfer of APE from large-scale to meso-scale

Transfer of KE from large-scale to meso-scale



Multi-Scale Dynamics

- Two distinct centers of instability: both of mixed type but different in cause.
- Center west of Pt. Sur: winds destabilize the ocean directly during upwelling.
- Center near the Bay: winds enter the balance on the large-scale window and release energy to the mesoscale window during relaxation.
- Monterey Bay is source region of perturbation and when the wind is relaxed, the generated mesoscale structures propagate northward along the coastline in a surface-intensified free mode of coastal trapped waves.
- Sub-mesoscale processes and their role in the overall large, mesoscale, sub-mesoscale dynamics are under study.



Energy transfer from meso-scale window to sub-mesoscale window.

Error Analyses and Optimal (Multi) Model Estimates Strategies For Multi-Model Adaptive Forecasting

- <u>Error Analyses</u>: Learn individual model forecast errors in an on-line fashion from model-data misfits based on Maximum-Likelihood
- <u>Model Fusion</u>: Combine models via Maximum-Likelihood based on the current estimates of their forecast errors

3-steps strategy, using model-data misfits and error parameter estimation

1. Select forecast error covariance ${\bf B}$ and bias ${\boldsymbol \mu}$ parameterization ${\boldsymbol \alpha}, {\boldsymbol \beta}$

$$\mathbf{B} pprox \tilde{\mathbf{B}}(\boldsymbol{\alpha}); \qquad \boldsymbol{\mu} pprox \tilde{\boldsymbol{\mu}}(\boldsymbol{\beta}); \qquad \boldsymbol{\Theta} = \{\boldsymbol{\alpha}, \boldsymbol{\beta}\}$$

2. Adaptively determine forecast error parameters from **model-data misfits** based on the Maximum-Likelihood principle:

$$\Theta^* = \arg \max_{\mathbf{\Theta}} p(\mathbf{y}|\mathbf{\Theta})$$
 Where $\mathbf{y} = \{\mathbf{y}_1^o, \mathbf{y}_2^o, \dots, \mathbf{y}_T^o\}$ is the observational data

3. Combine model forecasts via Maximum-Likelihood based on the current estimates of error parameters

O. Logoutov

Forecast Error Parameterization

Limited validation data motivates use of few free parameters

• Approximate forecast error covariances and biases as some parametric family, e.g. homogeneous covariance model:

$$\mathbf{B}_{m}(i,j) = \sigma(\mathbf{x}_{i})\sigma(\mathbf{x}_{j})\rho(||\mathbf{x}_{i} - \mathbf{x}_{j}||); \quad \rho(r) = \exp\left(\frac{-r^{2}}{2L^{2}}\right)$$

- Choice of covariance and bias models $\tilde{\mathbf{B}}$ and $\tilde{\boldsymbol{\mu}}$ should be sensible and efficient in terms of $\tilde{\mathbf{B}}\mathbf{v}$, $\tilde{\mathbf{B}}^{-1}\mathbf{v}$ and storage
 - * functional forms (positive semi-definite), e.g. isotropic
 - facilitates use of Recursive Filters and Toeplitz inversion
 - * feature model based
 - sensible with few parameters. Needs more research.
 - * based on dominant error subspaces
 - needs ensemble suite, complex implementation-wise

Error Parameter Tuning

Learn error parameters in an on-line fashion from model-data misfits based on Maximum-Likelihood

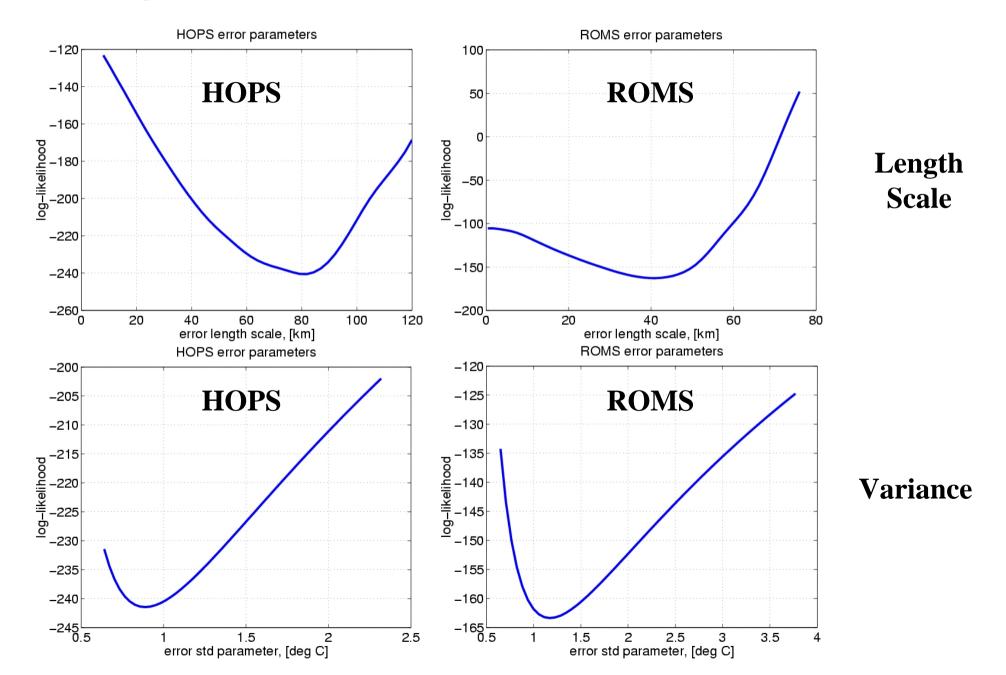
• We estimate error parameters via Maximum-Likelihood by solving the problem:

$$\mathbf{\Theta}^* = \arg\max_{\mathbf{\Theta}} p(\mathbf{y}|\mathbf{\Theta}) \tag{1}$$

Where $\mathbf{y} = \{\mathbf{y}_1^o, \mathbf{y}_2^o, \dots, \mathbf{y}_T^o\}$ is the observational data, $\Theta = \{\theta_1, \theta_2, \dots, \theta_M\}$ the forecast error covariance parameters of the M models

- (1) implies finding parameter values that maximize the probability of observing the data that was, in fact, observed
- By employing a randomized algorithm, we solve (1) relatively efficiently

Log-Likelihood functions for error parameters



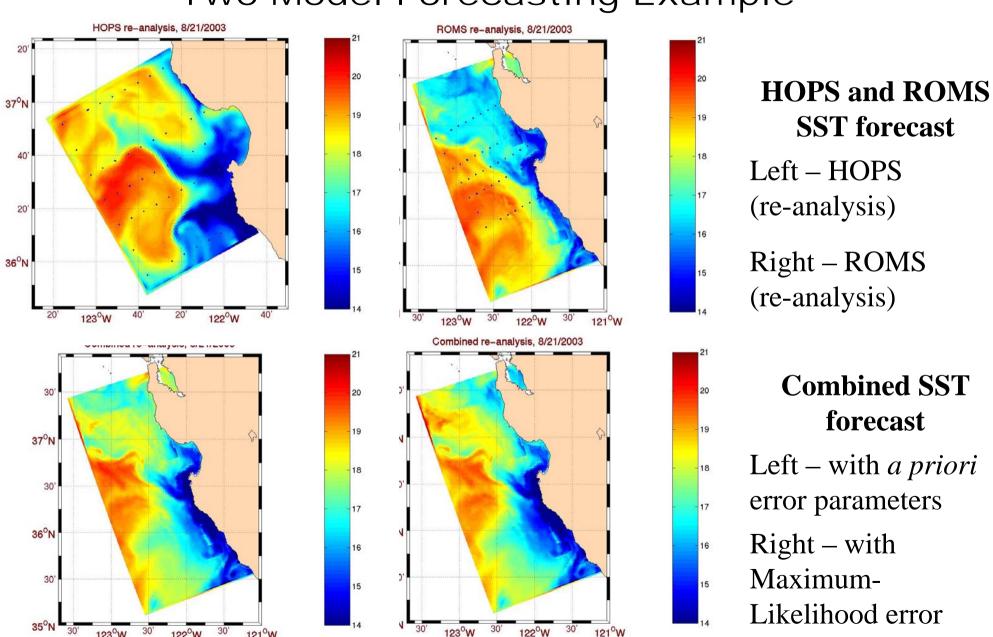
Model Fusion

combine based on relative model uncertainties

• Model Fusion: once error parameters Θ^* are available, combine forecasts \mathbf{x}_m based on their relative uncertainties as:

$$\mathbf{x}^* = \arg\min_{\mathbf{x}} \sum_{m=1}^{M} (\mathbf{x} - \mathbf{H}_m \mathbf{x}_m)^T \mathcal{B}_{(\mathbf{\Theta}_m)}^{-1} (\mathbf{x} - \mathbf{H}_m \mathbf{x}_m)$$

Two-Model Forecasting Example



parameters

AOSN-II Papers – In Preparation or Planned

- Real-Time Forecasting and Re-Analysis Fields A.R. Robinson, et al.
- Real-Time Error Prediction and Data Assimilation with ESSE P.F.J. Lermusiaux, *et al*.
- Multi-Scale Dynamical Processes, Mesoscale and Sub-Mesoscale Dynamics X.S. Liang and A.R. Robinson et al. (in draft form)
- Models Errors Parameter Estimation and Multi-Model Estimation –
 O. Logoutov and A.R. Robinson.
- A generalized biological model and application in Monterey Bay, California R. Tian, P.F.J. Lermusiaux *et al*.
- Adaptive Sampling in Oceanography and Meteorology: ESSE and ETKF. Lermusiaux and Majumdar
- Model Comparisons and Multi-Model Estimates (A.R. Robinson, Y. Chao, I. Schulman, *et al.*?)
- Data and model comparisons (authorship?)
- Comparison of observed and modeled biology S. Haddock, C. Herren, R. Tian, P.F.J. Lermusiaux, *et al.*

Monterey Bay / CCS - Planned Research

• Descriptive Dynamical and Statistical Interpretation of Reanalysis

Monterey Bay / CCS Research

- > Free surface HOPS reanalysis
 - Hierarchy of tidal forcing effects
- Domain revisions
 - Placement of southern boundary, Vertical discretization, Horizontal resolution
- > 2-way Nesting (including Temperature & Salinity based Feature Model)
- > Additional data: AUV, AXBT, CODAR, AVHRR, SSH
- ➤ Diurnal assimilation

Continue with Multi-Model Estimates

Continue Multi-Scale Energy and Vorticity Analyses

- ➤ Conclude MS-EVA study of relaxation event
- > Apply MS-EVA to onset and maintenance of upwelling
- ➤ Determine sampling requirements of full multi-scale fields to test dynamical processes predicted by MS-EVA

Continue with Balance of Terms Studies

- ➤ Primary Term by Term balances
- ➤ Heat, Salt and Momentum Flux balance;
- > Energy and Vorticity balances
- For each case: balances during onset, maintenance and transition of upwelling and relaxation events

Monterey Bay /CCS - Planned Research (Cont.)

• ESSE

- Quantitative evaluation of real-time uncertainty predictions
- Study of ESSE correlation and covariance functions (local, vertical average, along-shore average, etc)
- Evaluation of details of assimilation scheme
- Descriptive dynamical interpretation of uncertainty fields
- Quantitative OSSEs and coupled studies with Princeton, SIO, WHOI, Cal Tech, etc.

• Generalized Biological Modeling and Coupled Physical-Biological Simulations

- ➤ Improve biological initial conditions (dynamical field and parameter balances)
- > Check all parameters with S. Haddock and others; Finalize conversion of Fluorescence to Chl
- ➤ Higher-vertical resolution in upper-layers
- Continue to compare data and model; quantify biological forecast skill; improve biological model
- ➤ Assimilate biological data
- ➤ Descriptive dynamical interpretation of the simulated fields

EXTRA VUGRAFS